

# Lithospheric Heat Flow and Dynamics

---

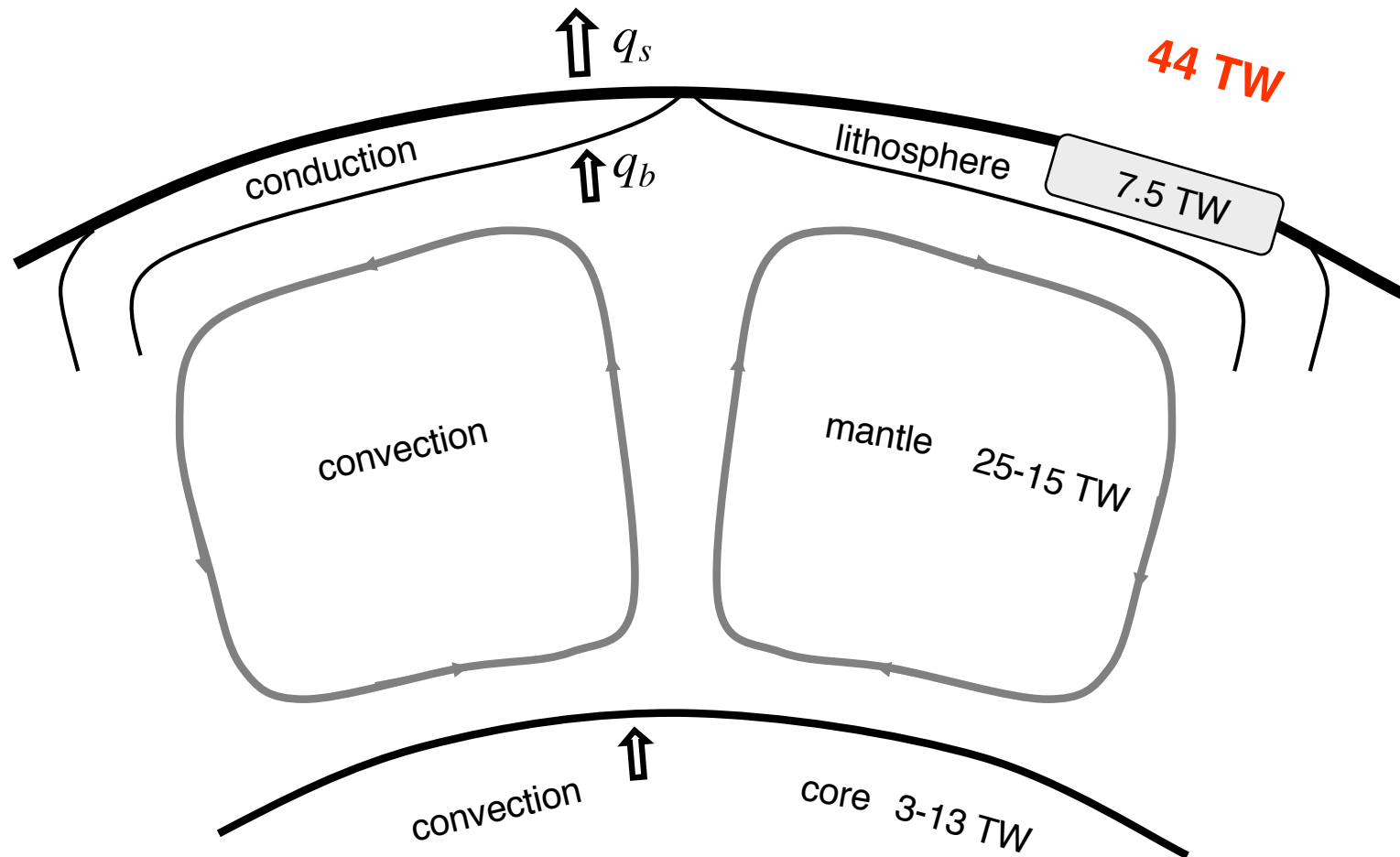
- **obvious signals**

- heat flow, depth, and geoid height versus age
- does hydrothermal circulation really transport 10 TW?

- **inferred signals**

- lithospheric thickness and strength versus age
- swell-push force and global stress from the geoid

# global heat budget

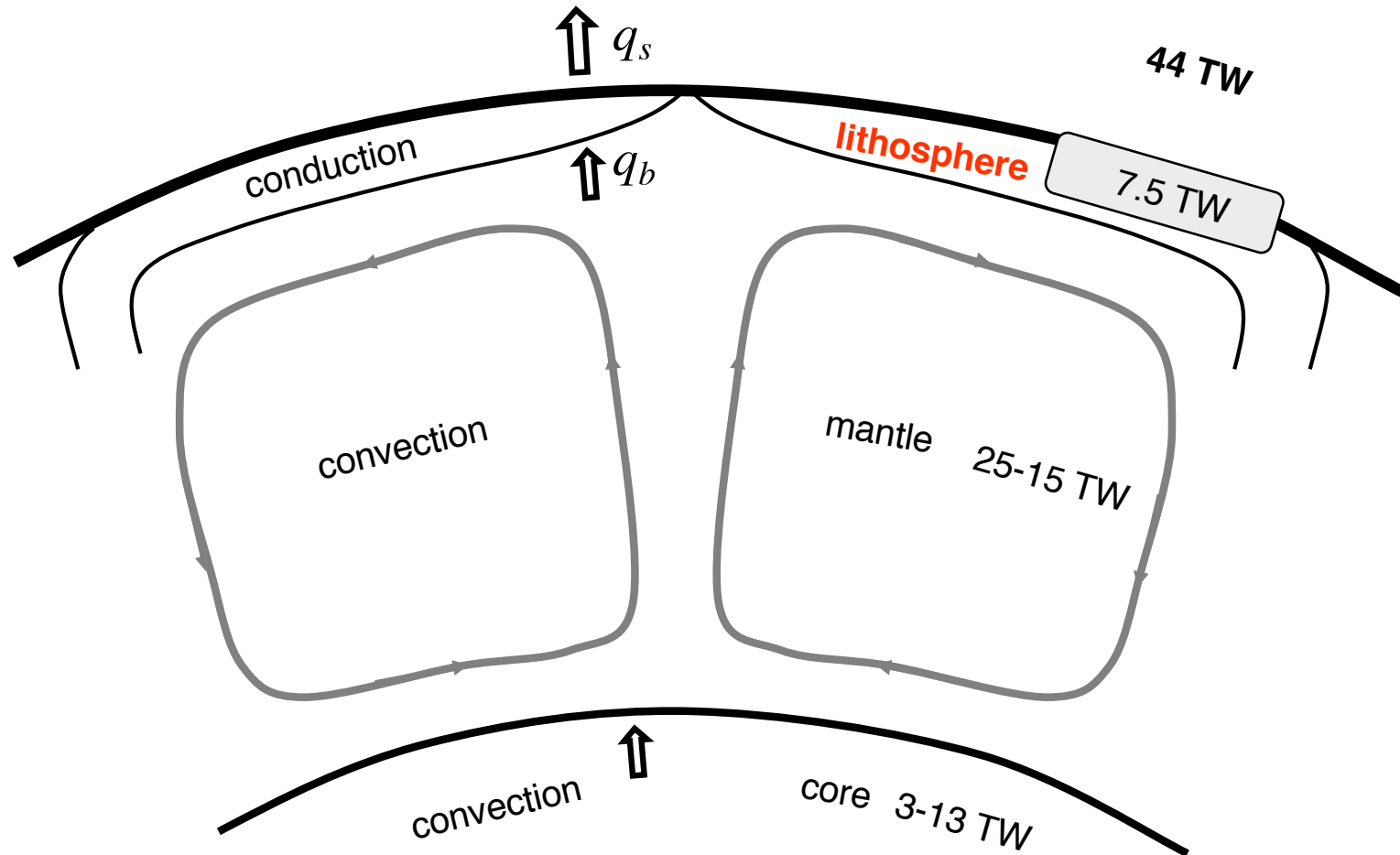


# oceanic lithosphere dominates mantle convection

largest surface area

greatest temperature drop across TBL = largest density contrast

> 1/2 of heat escapes in young oceanic lithosphere



ives  
he

tration C  
1)

$10^{-9}$

$10^{-9}$

$10^{-9}$

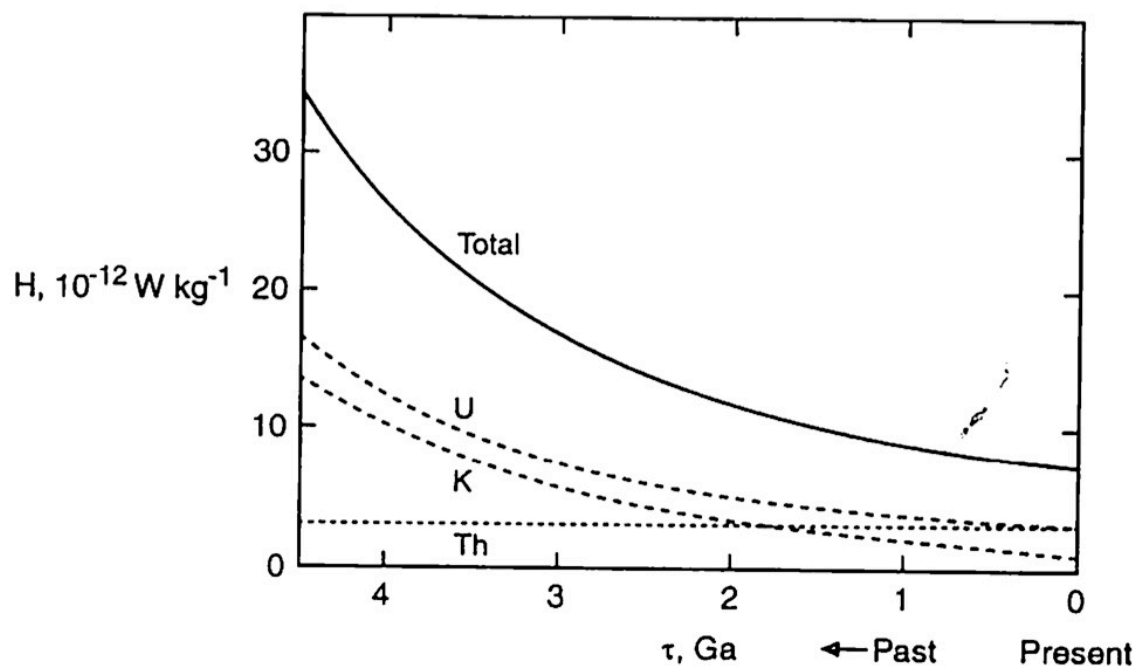
$10^{-9}$

$10^{-9}$

$10^{-5}$

oncentra-

ther para-  
find that  
er billion  
can man



**4-4** Mean mantle heat production rates due to the decay of the radioactive isotopes of U, Th, and K as functions of time measured back from the present.

plotted as a function of time before the present in Figure 4-4. The past contributions of the individual radioactive elements are also shown. We see that the rate of heat production  $3 \times 10^9$  yr ago was about twice the

1.5

1.4



# thermal expansion

---

volumetric expansion

$$\frac{\Delta V}{V} = \alpha \Delta T \quad \text{or} \quad \frac{\Delta \rho}{\rho} = -\alpha \Delta T$$

$\alpha$  - thermal expansion coefficient  $\sim 3 \times 10^{-5} \text{ } ^\circ\text{C}^{-1}$

linear expansion

$$\frac{\Delta l}{l} = \alpha_l \Delta T$$

$$\alpha_l \cong \frac{\alpha}{3}$$

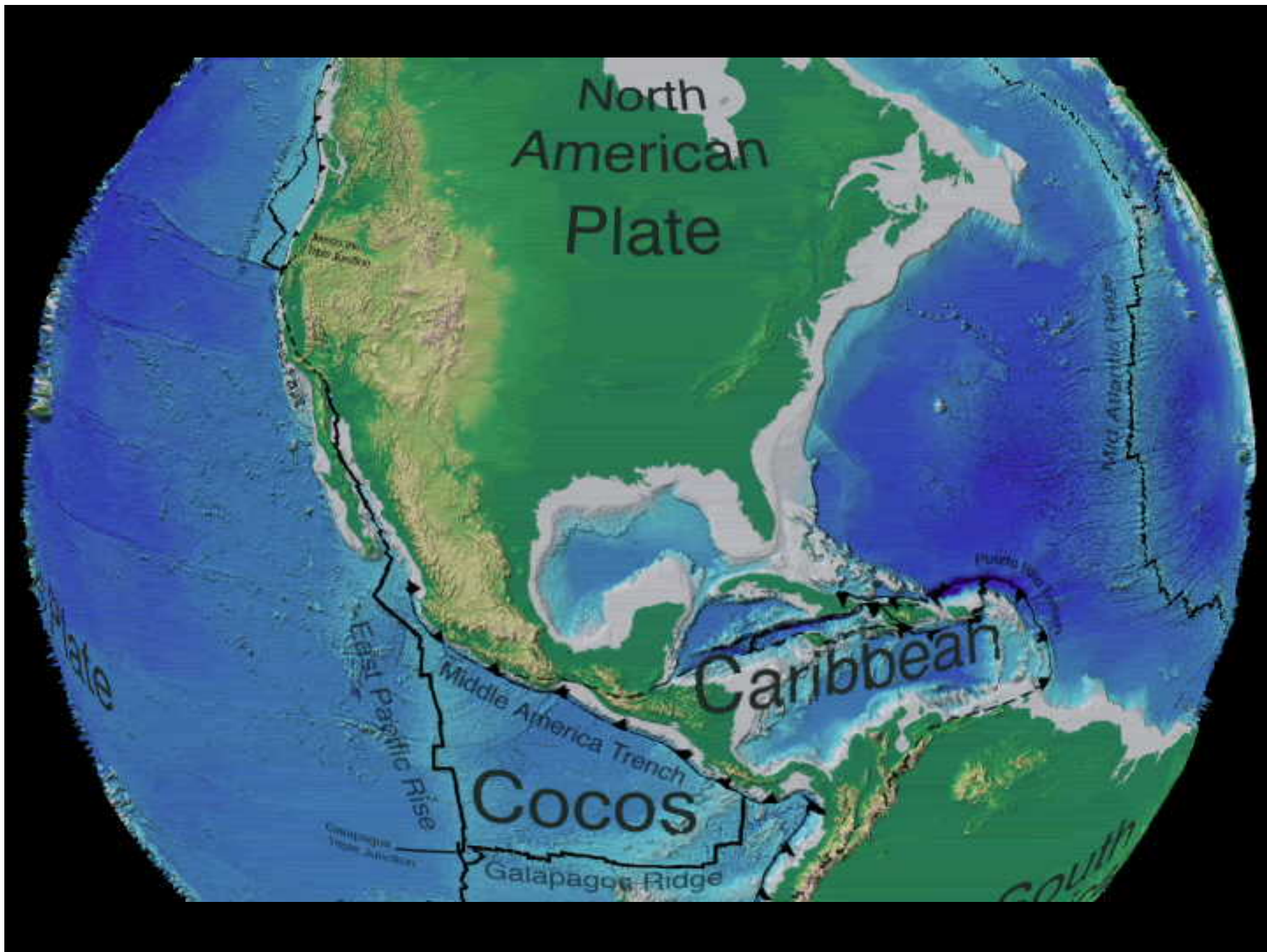
thermal stress  
develops when

$$\nabla(\Delta T) \neq 0$$

# obvious signals

---

- depth versus age
- heat flow versus age
- geoid height versus age



**depth vs age**  $\rightarrow d(t) = \frac{-\alpha \rho_m}{\rho_m - \rho_w} \int_0^L T dz \rightarrow d(t) \approx 2500 + 350t^{1/2}$

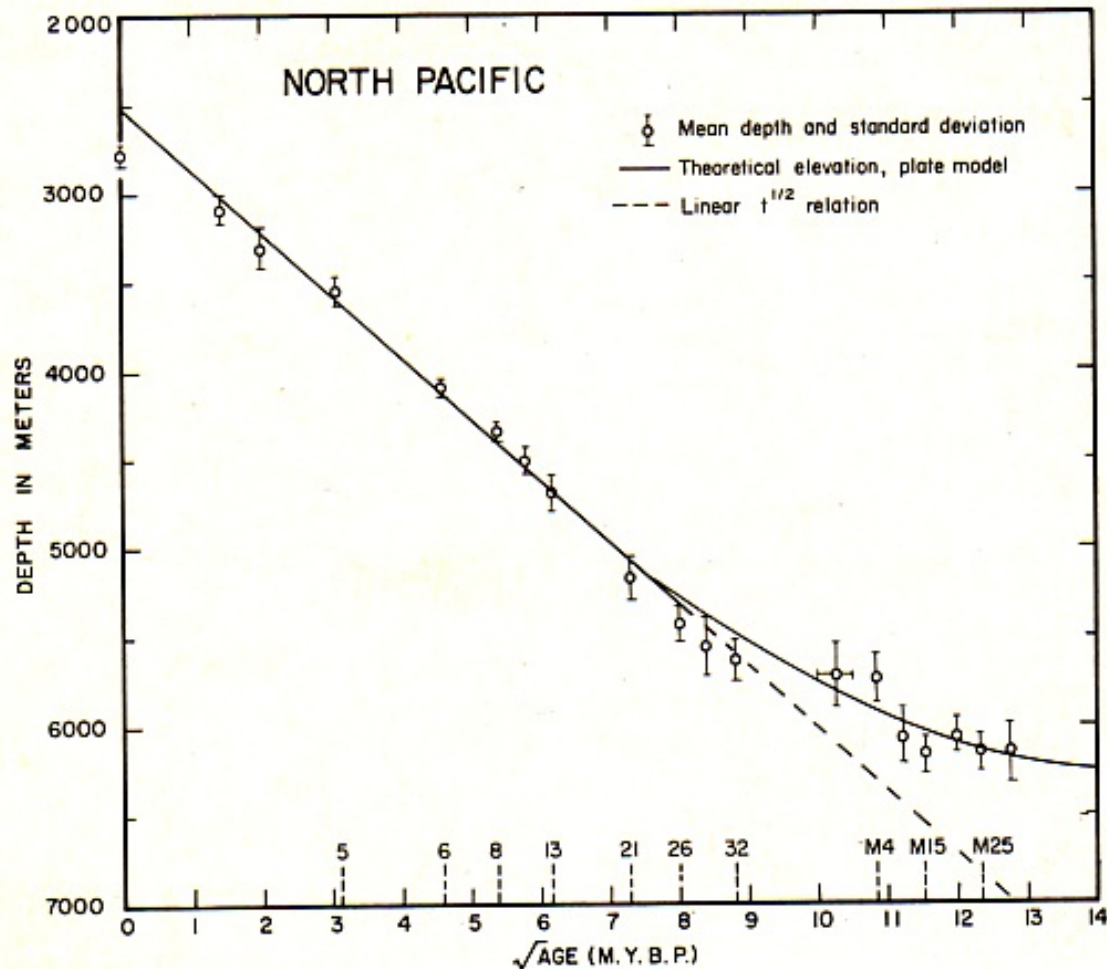
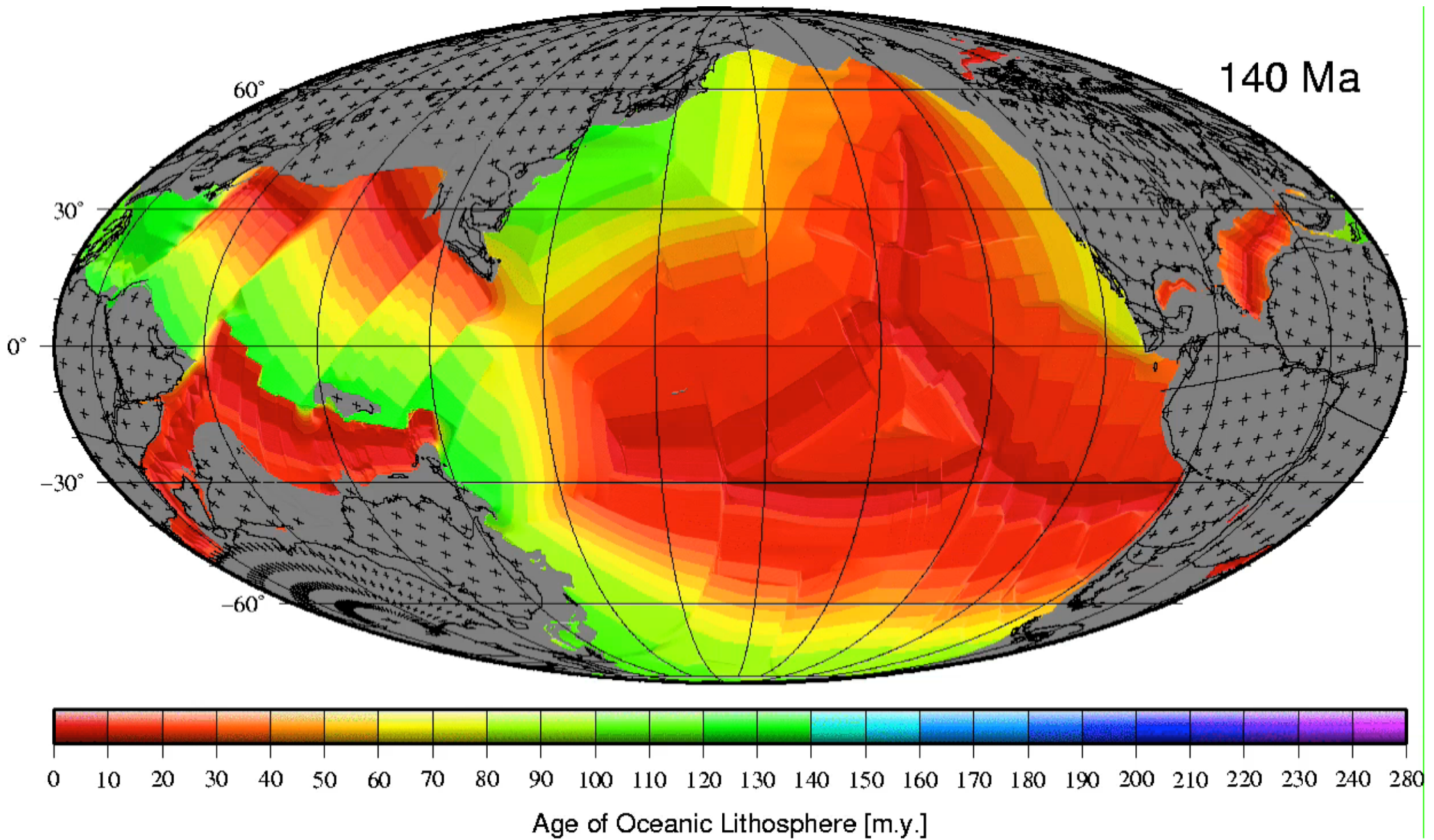
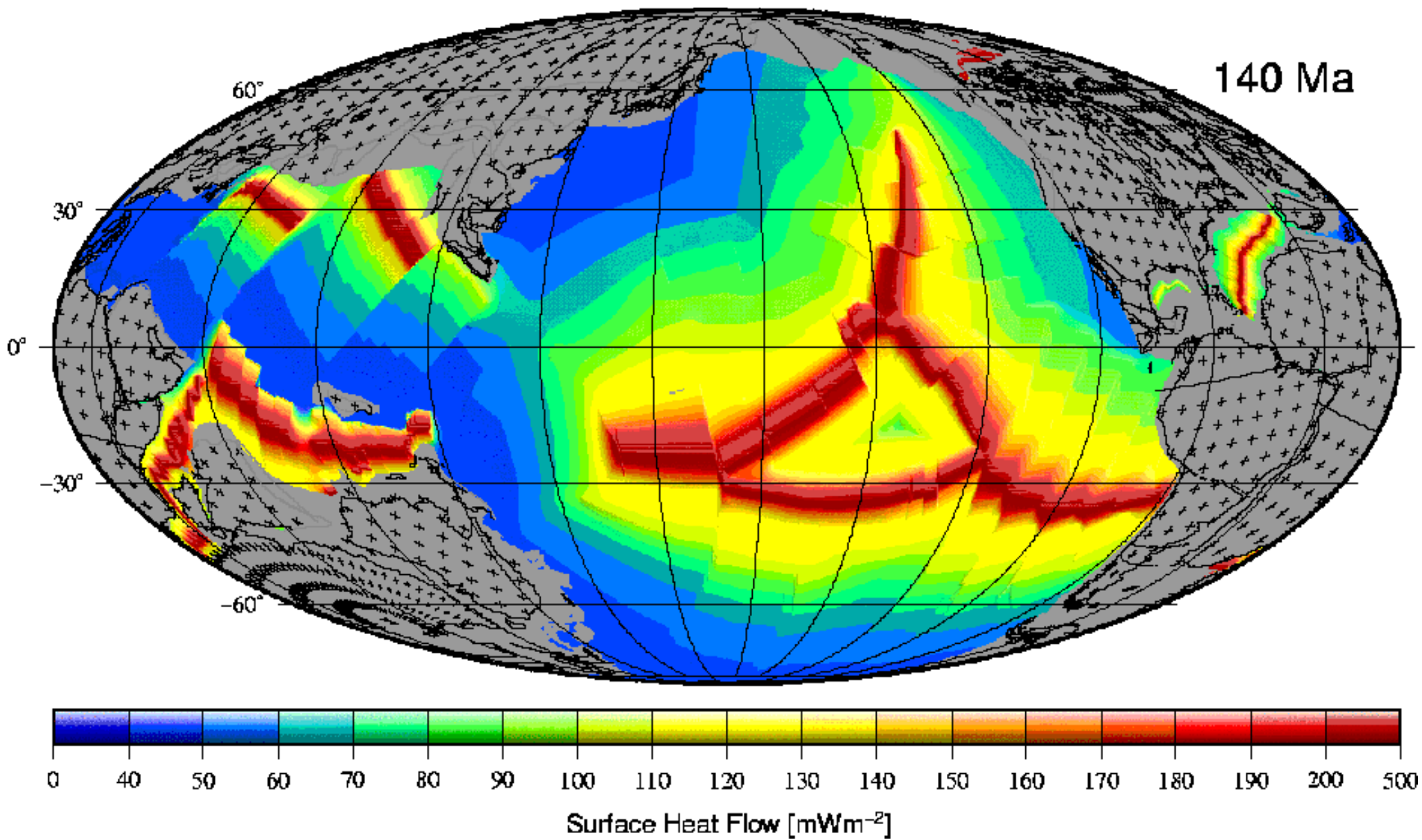


Fig. 1. Plot of mean depth in the North Pacific versus the square root of age. Numbers at the bottom of the figure denote selected Cenozoic and Mesozoic magnetic anomalies [from *Parsons and Sclater*, 1977].





Mueller, personal communication 2006



Mueller, personal communication 2006

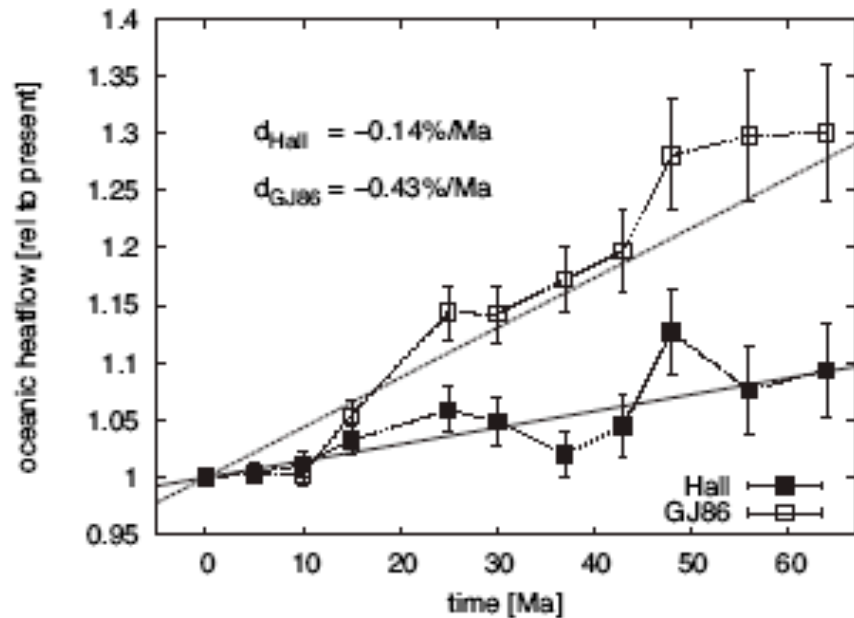


Fig. 2. Temporal variation of the globally integrated heat flow based on the estimates from the two plate tectonic reconstructions as shown in Fig. 1. For regions with half-space cooling-derived heat flow ( $q = C_A t^{-1/2}$ ), we performed an integration over seafloor age by summing over 1-Myr age integrals and multiplying each of these integrals by the area of seafloor within that age interval. Error bars for each heat flow estimate are computed by assuming progressively increasing uncertainty in these area estimates as described below. Filled square symbols (Hall) are for Xu et al.'s (6) reconstruction based on Hall (34), and open squares (GJ86) are based on Gordon and Jurdy (33). Solid and dashed lines are best-fit linear trends with rates,  $d$ , of relative change in total oceanic heat flow specified in the key.

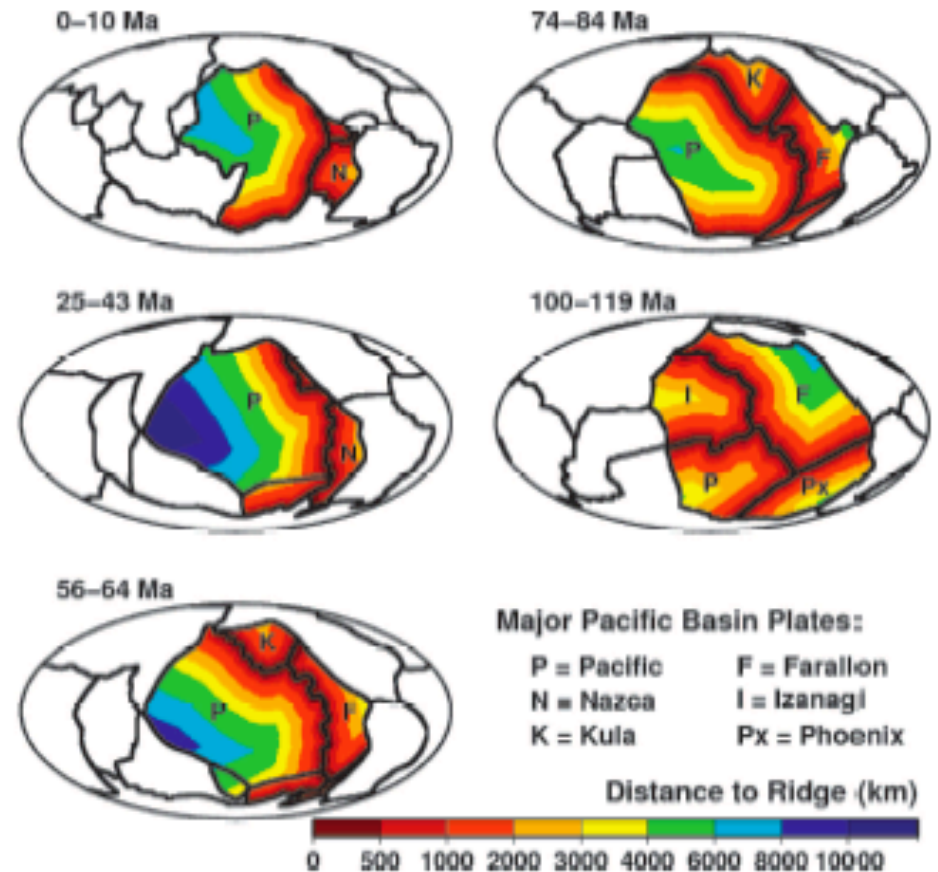


Fig. 3. Cenozoic and Mesozoic evolution of the Pacific basin as characterized by seafloor distance to the nearest mid-ocean ridge. Notice the progression from four relatively small plates to one large plate and compare with Fig. 1. Maps were created after Lithgow-Bertelloni and Richards (39).

Loyd, Becker, Conrad, Litho-Bertelloni and Corsetti, PNAS, 2007



# obvious signals - summary

---

## heat flow versus age

$$q_s(t) = k \frac{\partial T}{\partial z}$$

- surface temperature gradient
- noisy, observations  $\ll$  model

## depth versus age

$$d(t) = \frac{-\rho_m}{\rho_m - \rho_w} \int_0^L \alpha T dz$$

- integrated temperature

- **observations = model**

## geoid height versus age

$$N(t) = \frac{-2\pi G \rho_m}{g} \int_0^L \alpha T z dz$$

- first moment of temperature
- dominated by mantle geoid, observations  $\sim$  model



# Inferred signals

---

- lithospheric strength versus age (see *Watts*, 2001)
- swell-push force and global stress from the geoid

# Hawaiian-Emperor seamount chain

Plate  
Mechanics  
(flexure)

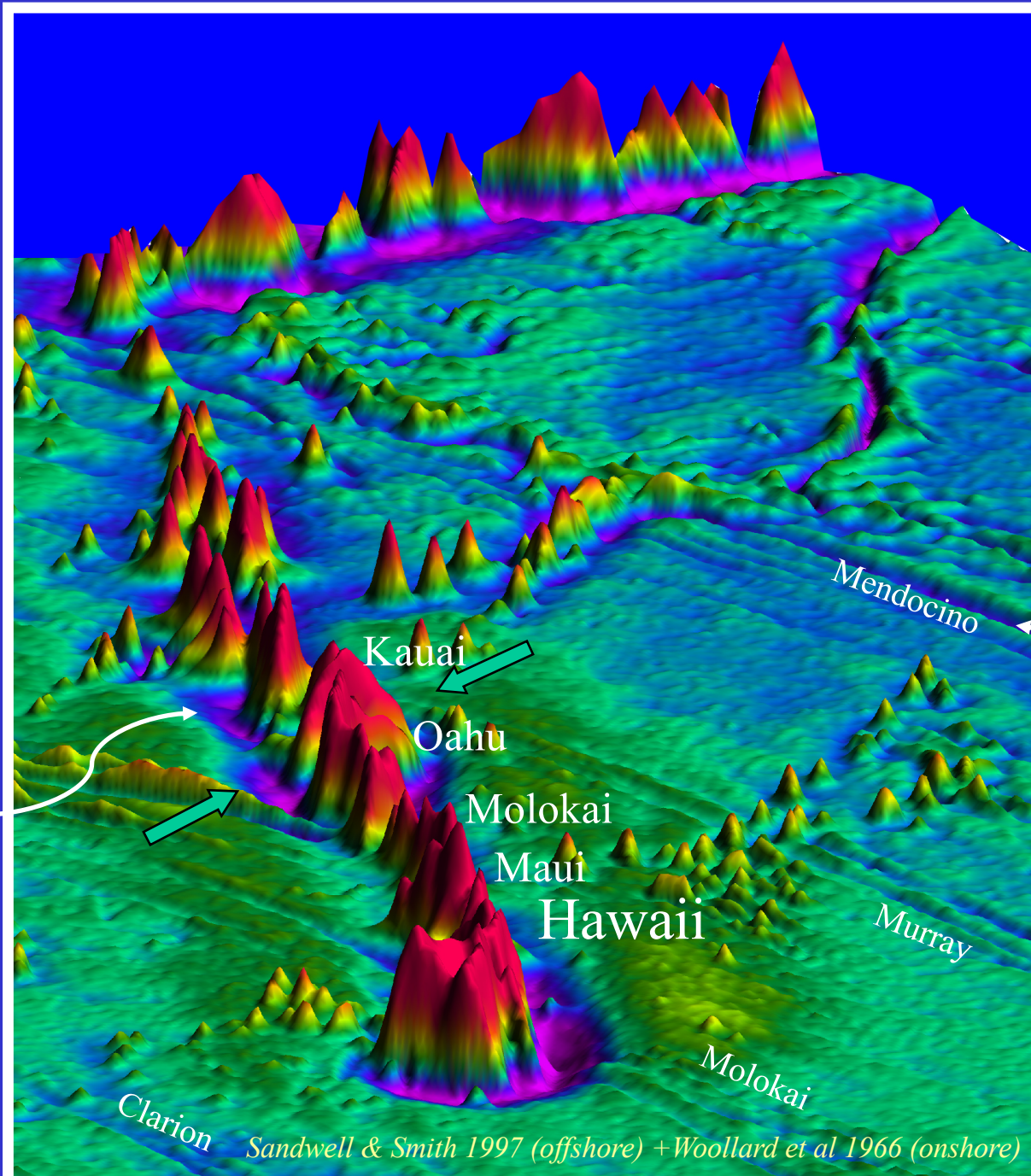
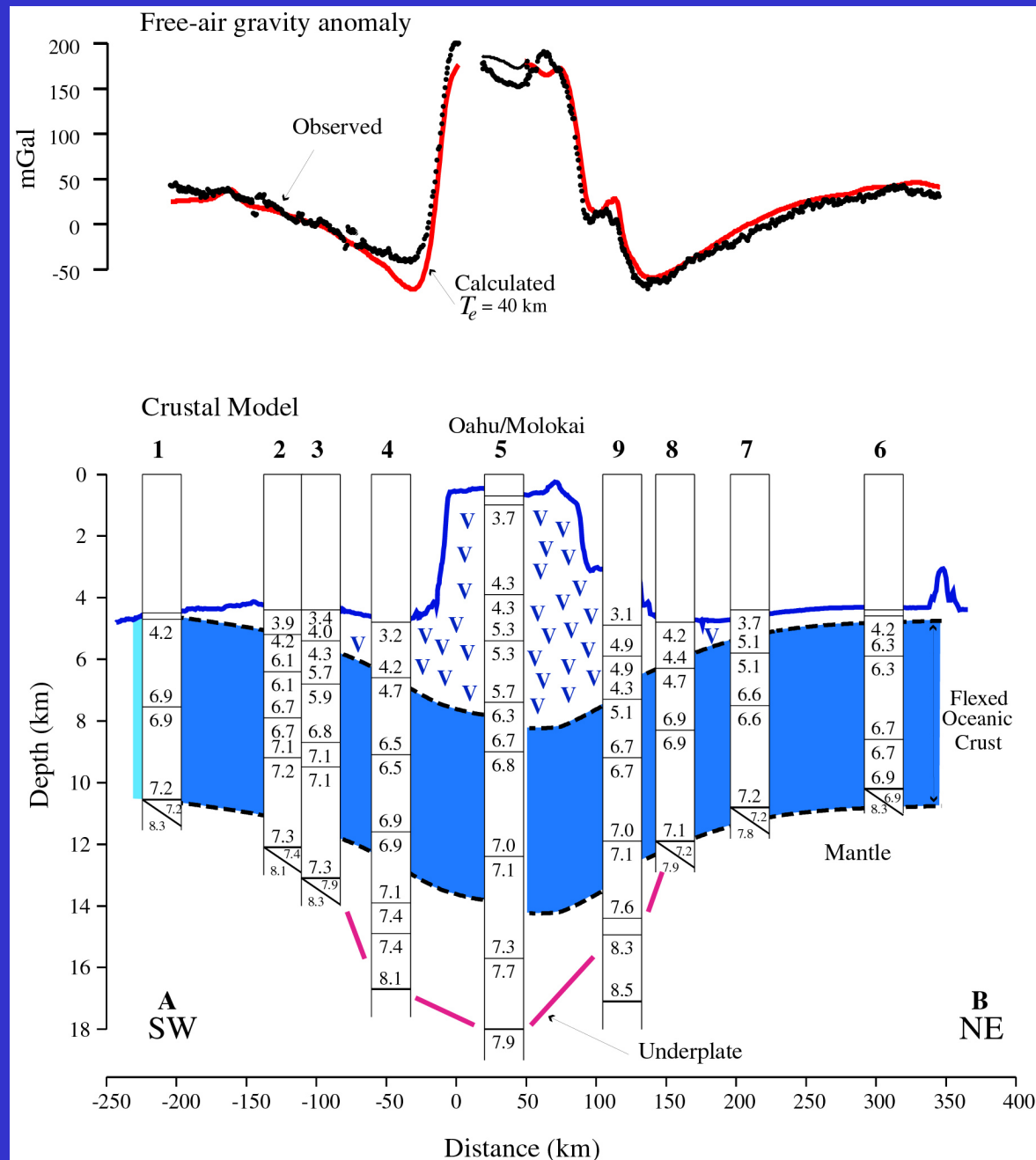
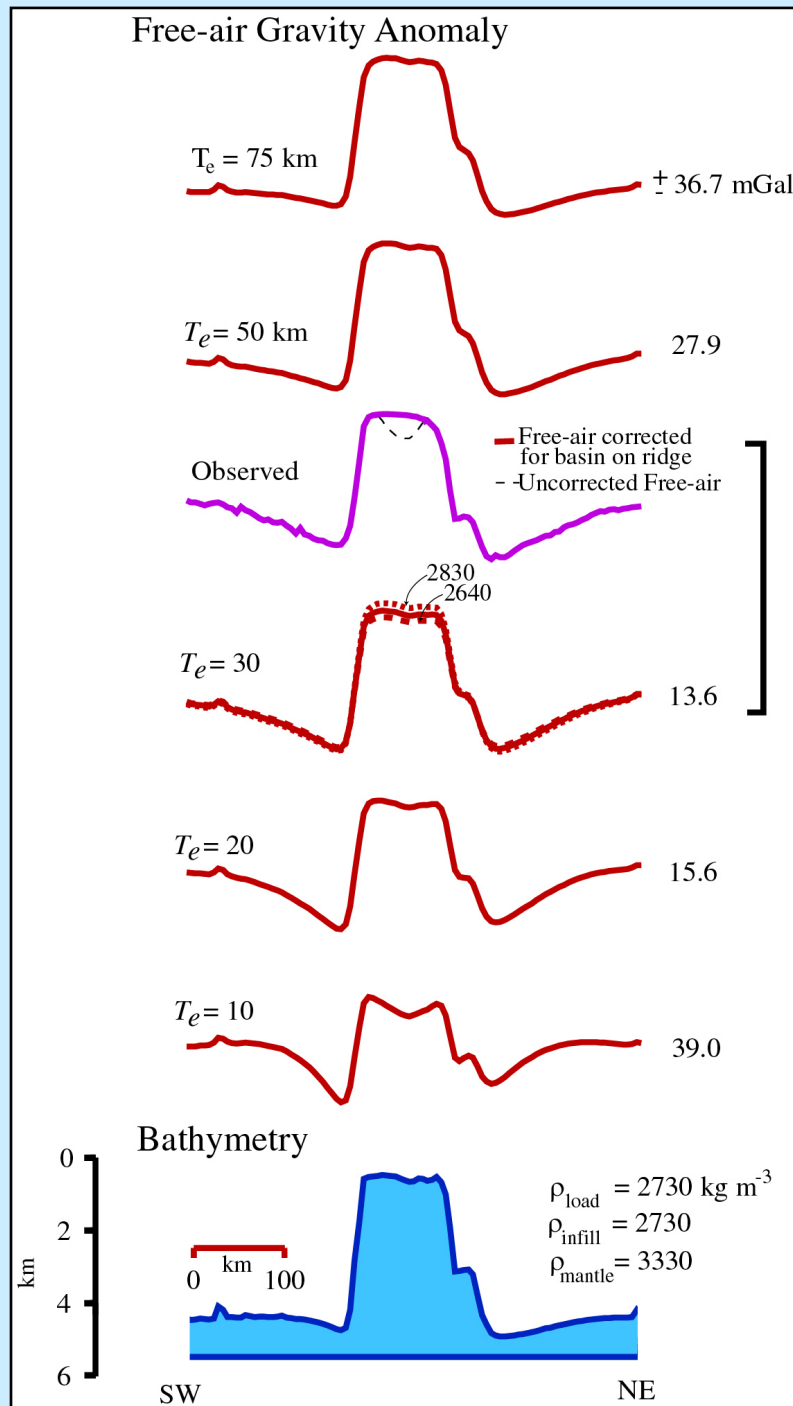


Plate  
kinematics

# Gravity anomalies and crustal structure at Oahu/Molokai

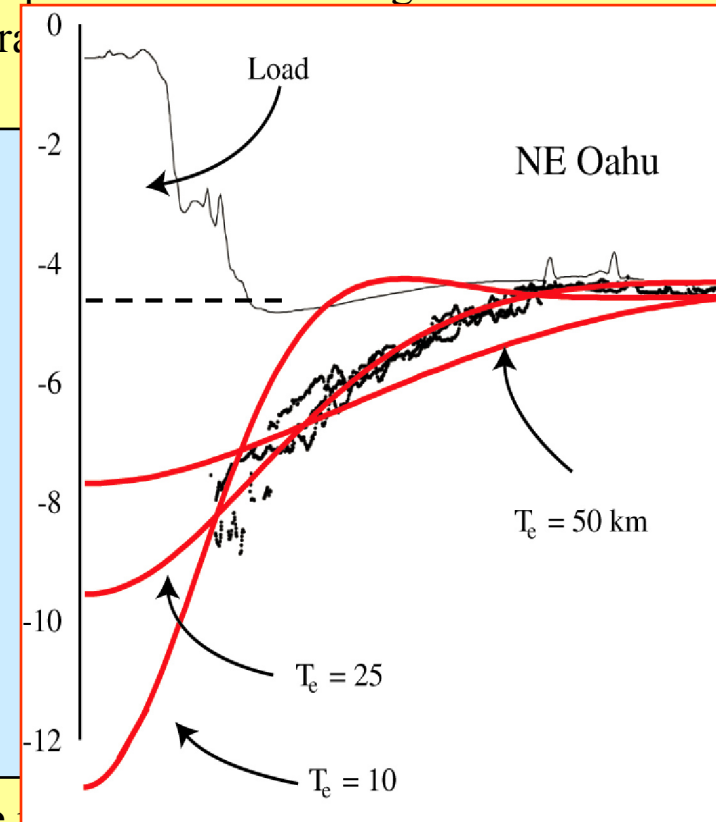


*Watts & ten Brink (1989)*



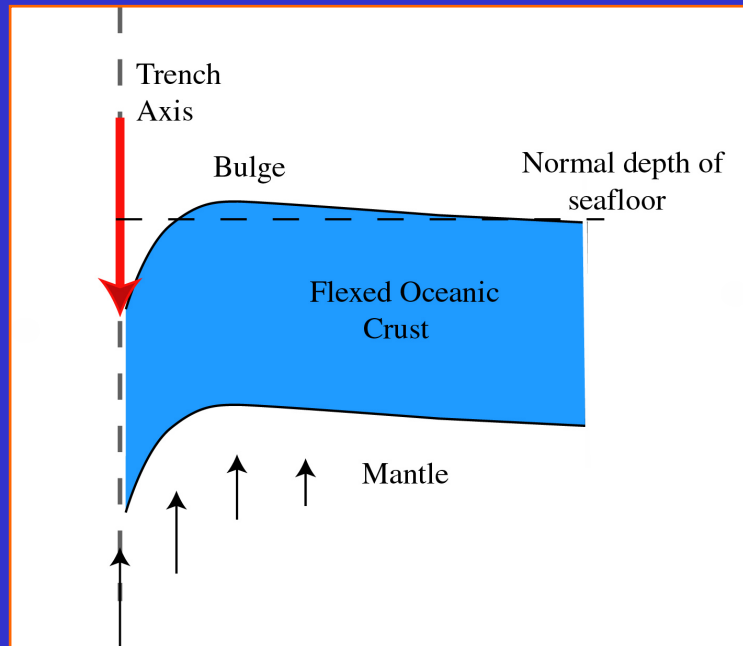
## Estimating $T_e$

$T_e$  can be estimated by comparing the amplitude and wavelength of the observed gravity anomaly with calculated curves for different  $T_e$  values.

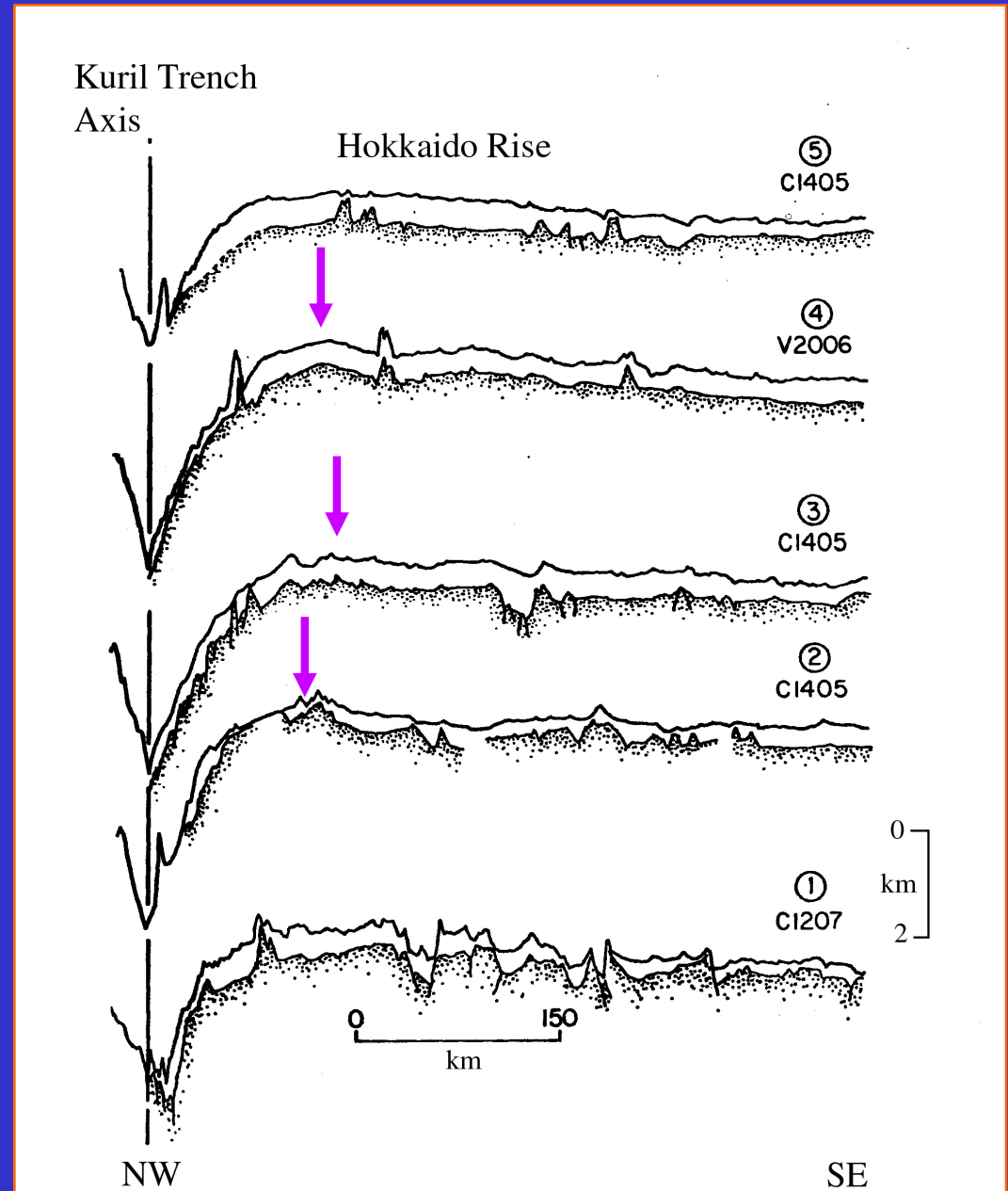


The minimum in the RMS difference between observed and calculated gravity anomaly indicate a 'best fit'  $T_e \sim 30 \text{ km}$ .

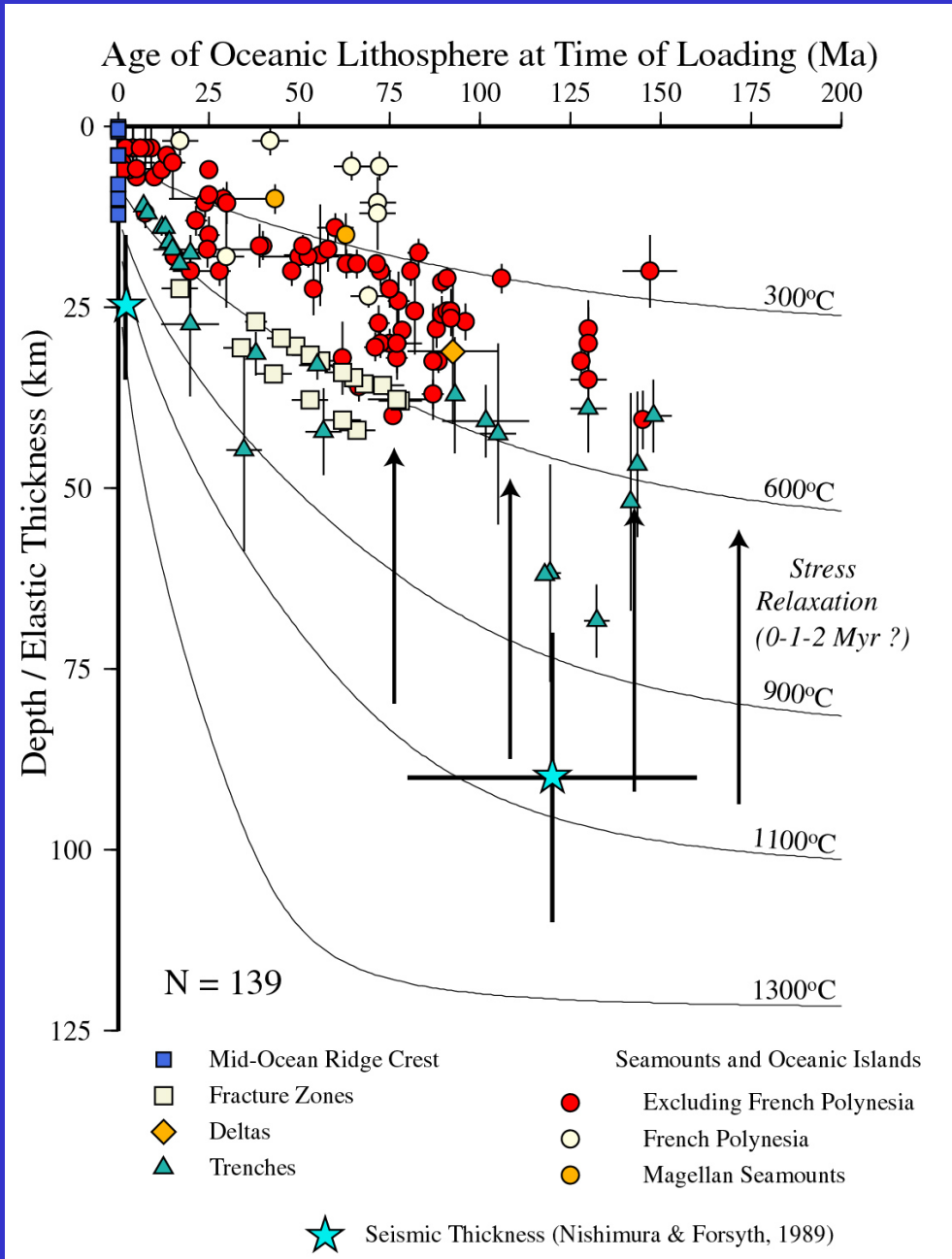
# Topography seaward of the Kuril Trench



Distance to bulge  $\sim 120\text{-}140$  km  
 $T_e \sim 30$  km



# Relationship between oceanic $T_e$ and plate and load age



# Lithospheric Heat Flow and Dynamics

---

- **obvious signals**

- heat flow, depth, and geoid height versus age
- does hydrothermal circulation really transport 10 TW?

- **inferred signals**

- lithospheric thickness and strength versus age
- swell-push force and global stress from the geoid

-