

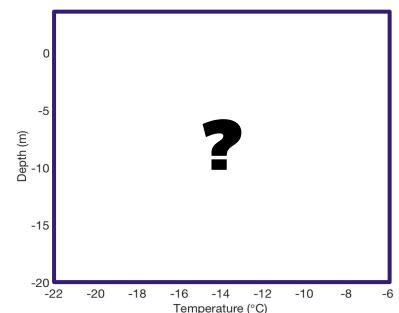
Nick Blanc & Margaret Morris

Problem

How does ice temperature close to a glacier surface vary with depth?

Things to consider:

- Seasonal temperature variations
- Heat diffusion dominates energy balance



Balance Equation

Energy Balance Equation → Advection-Diffusion Equation:

$$\rho C \left(\frac{\partial T}{\partial t} + \mathbf{v} \nabla T \right) = \nabla (k \nabla T) + P.$$
advection diffusion production

Simplified using constant k, and in 1-dimension:

$$\rho C \left(\frac{\partial T}{\partial t} + w \frac{dT}{dz} \right) = k \frac{d^2 T}{dz^2} + P.$$

Assumptions

Fourier law of heat diffusion

$$\frac{\partial T}{\partial t} = \kappa \frac{\partial^2 T}{\partial y^2}$$

We only care about one dimension (depth).

Assumptions

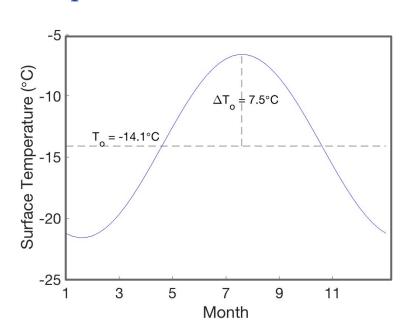
Boundary Conditions

 \circ Temperature changes seasonally with some amplitude ΔT_0

$$T(0,t) = T_0 + \Delta T_0 \sin(\omega t)$$

 \circ Constant at some depth (i.e. as $y \to \infty$)

$$T(\infty,t)=T_0$$



Separate and conquer!

$$T(y,t) = Y(y)R(t) \ = T_0 + Y_1(y)\cos(\omega t) + Y_2(y)\sin(\omega t).$$

$$\kappa rac{d^2Y_1}{dy^2} = \omega Y_2 \ ag{Guess!} \ \kappa rac{d^2Y_2}{dy^2} = -\omega Y_1 \ ag{Y_2} = C_0 e^{lpha y}$$

Assumed solution for Y₂

$$Y_2 = C_0 e^{\alpha y}$$

$$lpha^4 + \left(rac{\omega}{\kappa}
ight)^2 = 0 \qquad lpha = \pm \left(rac{1\pm i}{\sqrt{2}}\sqrt{rac{\omega}{\kappa}}
ight)^2$$

$$Y_1(y) = \expigg(-y\sqrt{rac{\omega}{2\kappa}}igg)igg[b_3\cos\Bigl(y\sqrt{rac{\omega}{2\kappa}} + b_4\sin\Bigl(y\sqrt{rac{\omega}{2\kappa}}\Bigr)igg]$$

$$Y_2(y) = \expigg(-y\sqrt{rac{\omega}{2\kappa}}igg)igg[b_1\cos\Bigl(y\sqrt{rac{\omega}{2\kappa}} + b_2\sin\Bigl(y\sqrt{rac{\omega}{2\kappa}}\Bigr)igg]$$

$$\kappa rac{d^2 Y_1}{du^2} = \omega Y_2$$

$$Ab_3\sin\left(y\sqrt{rac{\omega}{2\kappa}}
ight)-Ab_4\cos\left(y\sqrt{rac{\omega}{2\kappa}}
ight)=Ab_1\cos\left(y\sqrt{rac{\omega}{2\kappa}}
ight)+Ab_2\sin\left(y\sqrt{rac{\omega}{2\kappa}}
ight)$$

$$b_2 = b_3 = 0$$
 $b_1 = \Delta T_0 = -b_4$

$$T(y,t) = T_0 + Y_1 \cos(\omega t) + Y_2 \sin(\omega t)$$

$$T_0 = T_0 + \exp\left(-y\sqrt{rac{\omega}{2\kappa}}
ight)\cos(\omega t)igg[-\Delta T_0\sin\left(y\sqrt{rac{\omega}{2\kappa}}
ight)igg]$$

$$+\exp\left(-y\sqrt{rac{\omega}{2\kappa}}
ight)\sin(\omega t)igg[\Delta T_0\cos\left(y\sqrt{rac{\omega}{2\kappa}}
ight)igg]$$

$$T(y,t) = T_0 + \Delta T_0 \exp\left(-y\sqrt{rac{\omega}{2\kappa}}
ight) \sin\!\left(\omega t - y\sqrt{rac{\omega}{2\kappa}}
ight)$$

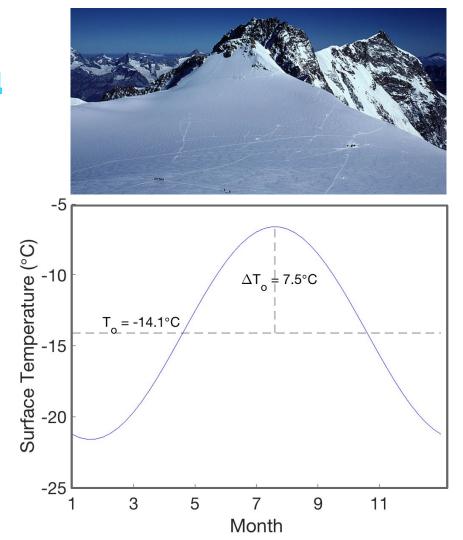
Glacier: Colle Gnifetti

Use borehole measurement temperatures from Haeberli & Funk 1991

Boundary Conditions:

$$T(0,t) = T_0 + \Delta T_0 \sin(\omega t)$$

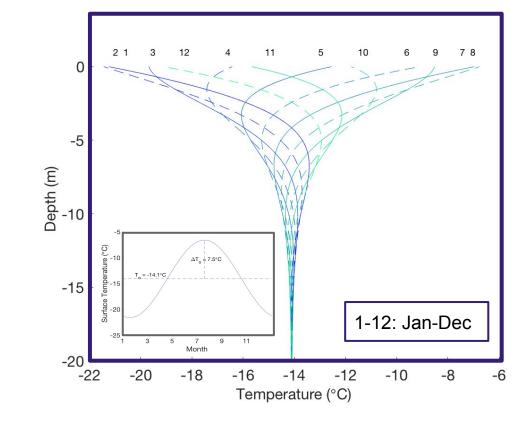
$$T(\infty,t)=T_0$$



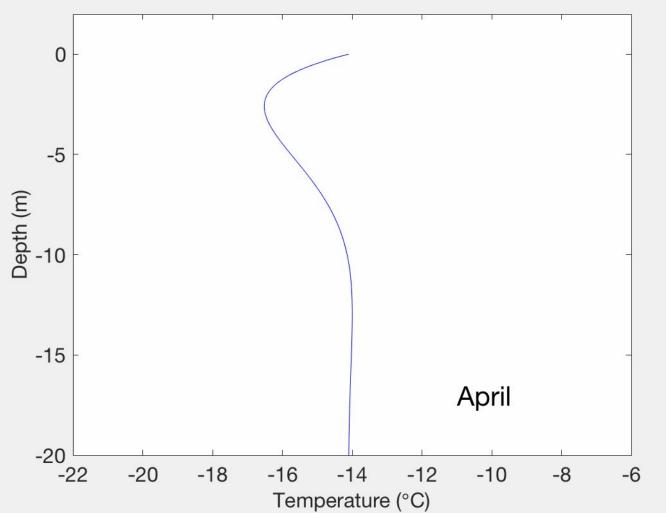
Result

Temperature varies with Depth

- Temp. variations small past 15m depth
- Temp settles to mean surface temperature at greater depths



$$T(y,t) = T_0 + \Delta T_0 \exp\left(-y\sqrt{rac{\omega}{2\kappa}}
ight) \sin\!\left(\omega t - y\sqrt{rac{\omega}{2\kappa}}
ight)$$

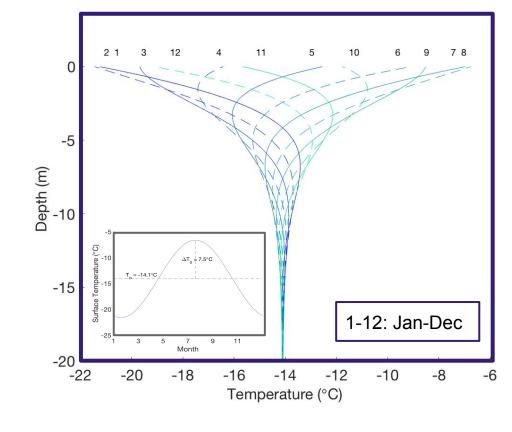




Summary

Deriving Temperature variation with Depth:

$$egin{aligned} rac{\partial T}{\partial t} &= \kappa rac{\partial^2 T}{\partial y^2} \ T(0,t) &= T_0 + \Delta T_0 \sin(\omega t) \ T(\infty,t) &= T_0 \end{aligned}$$



$$T(y,t) = T_0 + \Delta T_0 \exp\left(-y\sqrt{rac{\omega}{2\kappa}}
ight) \sin\!\left(\omega t - y\sqrt{rac{\omega}{2\kappa}}
ight)$$

References

https://www.igsoc.org/journal/37/125/igs journal vol37 issue125 pg37-46.pdf

Journal of Glaciology, Vol. 37, No. 125, 1991

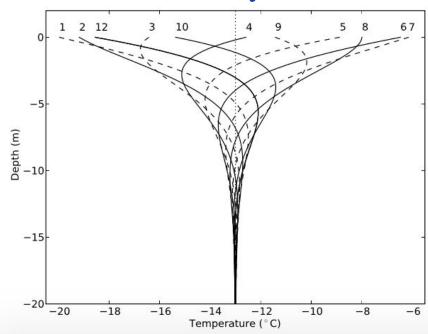
Borehole temperatures at the Colle Gnifetti core-drilling site (Monte Rosa, Swiss Alps)

WILFRIED HAEBERLI AND MARTIN FUNK
Versuchsanstalt für Wasserbau, Hydrologie und Glaziologie, ETH-Zentrum,
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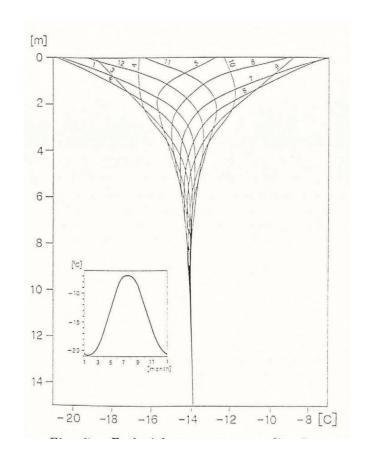
Cuffy Chapter 6

Extra Slides

Plot from Cuffy:

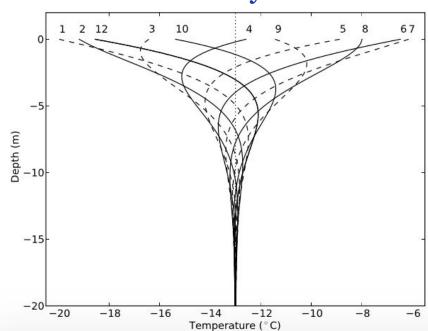


Plot from Haeberly & Funk:

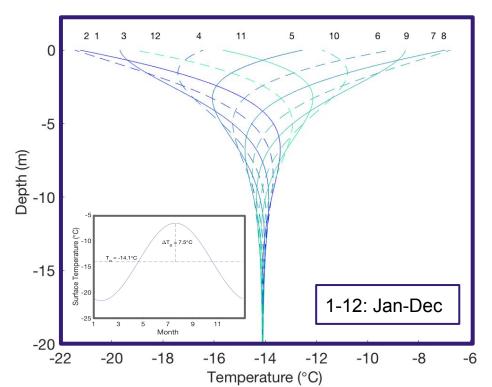


Extra Slides

Plot from Cuffy:



Plot we made:



Extra Slides

Plot from Haeberly & Funk:

Plot we made:

