CryoSat and Jason-1

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Abstract

Marine gravity anomalies derived from radar altimeter measurements of ocean surface slope are the primary data for investigating global tectonics and seafloor bathymetry. The accuracy of the global marine gravity field is limited by the availability of non-repeat altimeter data. Current models, having accuracies of 3-5 milligals (e.g., S&S V18 and DNSC08), are based on the non-repeat data collected by Geosat (18 mo.) and ERS-1 (12 mo.) which use altimeter technology from the 70's and 80's, respectively. The next opportunities for significant improvements in marine gravity will come from CryoSat and perhaps Jason-1 if it is placed into a non-repeat orbit. In addition to complete ocean coverage, the three attributes needed for improved gravity are improved range precision, optimal satellite inclination, and long mission duration. Range precision is established through a coherence analysis of repeating profiles. We will use these standard methods to assess the range precisions of CryoSat and Jason-1 in relation to Geosat, ERS, Topex, and Envisat. We expect that the higher PRF's of the newer altimeters will provide a square root of 2 improvement in range precision. The low inclination of the Jason-1 orbit offers the best opportunity for improvement in the E-W variations in gravity, especially at low latitude. The hopefully, long mission duration of CryoSat provides the best opportunity for reducing the noise due to ocean surface roughness from swells. Data from these two missions may eventually result in a factor of two improvement in the accuracy of the marine gravity field and geoid at scales less than 100 km.







Factor of 2 Improvement in Marine Gravity from

(a) Average of 27,000 CryoSat radar waveforms (solid) and a simplified model with four adjustable parameters: Aamplitude, t_{o} – arrival time, σ – rise time, and κ – trailing edge decay. The spacing of the gates is 3.12 nsec or 468

CryoSat

 $P(t,t_o,\sigma,A) = \frac{A}{2} \left\{ 1 + erf\left[\frac{(t-t_o)}{\sqrt{2}\sigma}\right] \right\} \quad \begin{cases} 1 & t < t_o \\ \exp[-\kappa(t-t_o)] & t \ge t_o \end{cases}$

(b) Significant wave height (swh) derived from 4-parameter retracking model (blue curve) has unrealistic short wavelength jitter. Smoothed swh (black curve) was provided to a 2-parameter retracker to constrain the shape of the leading edge of the waveform so the arrival time is more accurately estimated.

(c) Arrival time converted to along-track slope (black curve) compared with slope from the EGM2008 model (blue curve, essentially V18.1 gravity). Both were low-pass filtered at 18 km wavelength for direct comparison with a previous analyses of Geosat and ERS-1 data. The slope differences have a median absolute deviation of 2.46 microradian – part of this difference is gravity signal and part is altimeter noise.

(d) CryoSat waveforms used for the analysis. The anomaly at the equator is perhaps due to a rain cell. Further algorithm development is needed to identify and remove these bad data.



profile	# good	# bad	3-parameter retrack		2-parameter retrack	
			median (µrad)	mad (μ <mark>rad</mark>)	median (µrad)	mad (μ <u>rad</u>)
1	20722	327	0.0757	3.0352	0.0120	2.4649
2	20743	305	0.0208	3.3795	-0.0041	2.6484
3	20960	88	-0.1472	2.7753	-0.1556	2.1928
4	20963	83	-0.0240	2.7438	0.0172	2.2733
5	20737	307	-0.0763	3.1039	-0.0687	2.4448
6	20851	191	-0.0131	3.3554	0.0531	2.5796
7	20928	111	-0.0417	3.4093	-0.0313	2.6628
8	21002	29	0.0699	2.6920	0.0141	2.1061
tot	166906	1144		3.0618		2.4216
ERS-1				3.5633		3.3133
Geosat				4.576		3.4270

2-parameter retrack provides 21% improvement with respect to 3parameter retrack

ratio of ERS to Cryosat 1.3682 ratio of Geosat to Cryosat 1.4152